Physical and Numerical Analysis of Hurricanes

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Challenge

Goal: Assign a probability that a target region, for which we have no historical data, would be hit by a Tropical cyclone.

Data:

- Track in Latitude/Longitude
- Max wind speed
- Central Pressure

Issue:

- Probability of exact hypothetical track is extremely small
- Hurricane dynamics are complex and not fully understood

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Overview of method - Partitioning approach

How to estimate the Hurricane trajectory?

$$\mathbf{v} = \frac{d\mathbf{x}}{dt} = \mathbf{c} + \mathbf{U} \tag{1}$$

- c: Advection due to Coriolis Force. Solve vorticity equation in absence of background flow.
- U: Advection due to large-scale atmospheric wind flow. Fit each hurricane track with an ellipse function and obtain an empirical probability distribution F(u).

Inverse Problem: $\mathbf{v}, \mathbf{c} \rightarrow \mathbf{U} \rightarrow F^{-1}(\mathbf{U})$

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Vorticity equation

We consider a 2D Navier-Stokes equation in a barometric framework.

$$\frac{\partial \zeta}{\partial t} + u \frac{\partial \zeta}{\partial x} + v \frac{\partial \zeta}{\partial y} + \beta v = 0$$

u = (u, v) velocity vector of hurricane
 β = df/dy Derivative of Coriolis parameter

• $\zeta = \mathbf{k} \cdot (\nabla \times \mathbf{u})$ relative vorticity

Initial State: symmetric vortex at the origin with an imposed tangential and angular velocity vector V(r), $\Omega(r) = V(r)/r$.

Tangential and Angular velocity



Over the time, the hurricane develops asymmetries due to the interaction with the ambient flow that generates itself.

Partitioning problem

We write $\zeta = \zeta_s + \Gamma$ to split ζ into an axisymmetric part for the core of the hurricane which just rotates, and an asymmetric correction term Γ . After some simplifications we can then split the vorticity eqn into two pieces:

$$\frac{\partial \zeta_s}{\partial t} + \mathbf{c}(t) \cdot \nabla \zeta_s = 0 \tag{3}$$

$$\frac{\partial \Gamma}{\partial t} = -\mathbf{u} \cdot \nabla (\Gamma + f) \tag{4}$$

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Streamfunction $\boldsymbol{\Psi}$

In cylindrical coordinates (r, θ) , the equation for Γ is written as:

$$\left[\frac{\partial}{\partial t} + \Omega(r)\frac{\partial}{\partial r}\right](\Gamma + \beta y) = 0$$
(5)

From $\Gamma,$ we can solve the Poisson equation $\nabla^2\Psi=\Gamma$ to obtain the Streamfunction. It follows that:

$$\Psi = \Psi_1(r, t) \cos(\theta) + \Psi_2(r, t) \sin(\theta)$$
(6)

The Streamfunction contour levels represent the direction along which the vortex of the hurricane moves.

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Solution for vortex speed

Finally the vortex speed is given by:

$$\left[X(t), Y(t)\right]^{T} = \begin{bmatrix} \frac{1}{2}\beta \int_{0}^{\infty} r\left[t - \frac{\sin\{\Omega(r)t\}}{\Omega(r)}\right] dr\\ \frac{1}{2}\beta \int_{0}^{\infty} r\left[1 - \frac{1 - \cos\{\Omega(r)t\}}{\Omega(r)}\right] dr \end{bmatrix}$$
(7)



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In addition to the component c, the other cause of the large scale hurricane motion is the environment velocity U.

We compute different U by fitting the Hurricane tracks provided by the Historical data to an ellipse. We also obtain an empirical distribution F(U) to sample from.



Inverse Problem

If we have an idea of a potential hurricane speed that might hit a region, we can extract the environment field U from eq. 1 and obtain the corresponding probability from its CDF.



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Summary

- Geophysics suggests that hurricanes move due to two distinct mechanisms.
- The interaction between the hurricane vortex and the Coriolis force term can be estimated analytically, while large-scale atmospheric flows additionally steer it around the Atlantic basin.

For future work:

- Combine the two elements to produce a semi-analytic model that estimates whole tracks
- We should find more regularity in the distribution of these parameters that leads to more reliable estimates for as-yet unseen tracks

Thank you for your attention!



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Solution for Environment Vorticity Equation

The equation 5 is integrated to give the solution:

$$\Gamma(r,\theta,t) = \zeta_1(r,t)\cos(\theta) + \zeta_2(r,t)\sin(\theta)$$
(8)
where $\zeta_1 = -\beta r \sin\Omega(r)t$ and $\zeta_2 = -\beta r [1 - \cos\Omega(r)t].$

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